



Agents of Physical Weathering









Biological Activity

Roots extend into cracks and expand the cracks as the roots grow in diameter.

 Biological activity also promotes chemical weathering.

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Chemical Weathering

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NOTE

- Many minerals are formed at conditions of temperature and pressure that are much different from those at the surface of the earth. To them, the earth's surface is a cold vacuum with destructive chemicals (oxygen, water, etc).
- Thus, many minerals begin to change into newer minerals that are in equilibrium with the surface temperatures and pressures. Rarely do the new minerals occupy the same volume or space as the original minerals. When a feldspar changes to clay, the clay takes up more space.
- The bulk chemical compositions may stay the same but the arrangement of atoms changes (e.g. minerals change).
- Many of the minerals in soils form because of chemical weathering.

Chemical weathering is most important in areas with lots of liquid water and warm temperatures, i.e. the tropics.

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- Increasing surface area with increasing disaggregation
- Weathering is more extreme with higher temperature minerals. Olivine weathers much, much faster than quartz.







| Hydrolysis |
|------------|
|------------|

| TABLE 6.1 | Products of Weathering | |
|---|---|---|
| Mineral | Residual Products | Material in Solution |
| Quartz Feldspars Amphibole (hornblende) Olivine | Quartz grains Clay minerals Clay minerals Limonite Hematite Limonite Hematite | Silica Silica K ⁺ , Na ⁺ , Ca ²⁺ Silica Ca ²⁺ , Mg ²⁺ Silica Mg ²⁺ |
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Factors Governing Soil Formation

- 1. Climate soils develop faster in areas with warm and humid climates.
- Parent Material usually either residual or transported materials. Soils form faster with residual parent material.
- 3. Slope greater slope increases runoff and therefore decreases the time for water to infiltrate. Soils form better with low slopes. The best soils form in areas that are slightly undulating and well drained.
- 4. Time usually the longer a soil has been developing the thicker and better developed the soil is.
- Biologic Activity since organisms furnish organic matter for soils, the more biologic activity there is the better developed the soils will be.

Soil Profile O horizon Loose and parity decayer organic matter A horizon Minesal matter mixed with some humus E horizon Light colored mineral partici Zone of eluxiation and leaching O Horizon - loose and partly. decayed organic material. A Horizon - Mineral matter mixed with some humus B horizon Accumulatio clay transportion above • E Horizon - Zone of eluviation and leaching C horizon Partially altered parent material B Horizon - Accumulation of transported clay from above. C Horizon - Partially altered parent Unweathered parent material material. Parent material Copyright @ 2005

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TYPES OF SOILS

PEDALFERS - Mainly in eastern US. Name refers to the aluminum-rich clay and iron oxides in the red to brownish B horizon. They are clayey soils that develop in areas with plentiful rainfall.

- PEDOCALS Mainly western US, in areas with little water, not very much clay because there isn't much chemical weathering. Leaching is slower and less clay is produced. Insufficient ground water to flush the calcium carbonate from the A horizon so it accumulates there as water evaporates. Can form caliche layers.
- LATERITES soils that form in the hot and humid tropics. Very thin organic layer over a very thick leached layer. Weathering is so intense that all minerals are decayed - even quartz is removed. This leaves a residue rich in aluminium and iron oxides and the soils are very red. Tropical soils have so little organic material that when forests are stripped and agriculture begins the soil is exhausted in just a few years. There be much vegetation, but organic material is rapidly decomposed by bacterial action. Ancient laterities form deposits called bauxite that are the principal aluminum ores of the world

